

Preliminary AURIC & libRADtran background simulation for EUSO-Balloon flight

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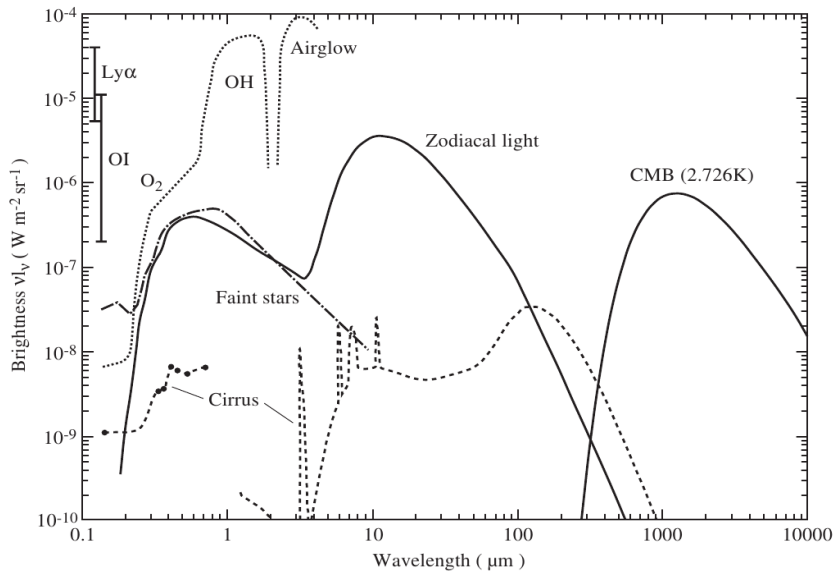
Part I: Sources of background for balloon

- To simulate situation for EUSO-balloon we extract:
 - Zodiacal light spectrum from Leinert [1]
 - Star light spectrum from Leinert [1]
 - Airglow spectrum from AURIC model [2]
 - These spectra was propagate through atmosphere by libRADtran to simulate balloon measurement at 40 km. [3]
 - Surface mixed forest – albedo 0.028 (water albedo doesn't change situation see my Moscow presentation)
 - No moon, no city lights, no aurora, cloudless
 - Preliminary value of background from EUSO-balloon data for cloudless condition at ~ **5:30 UT is ~ 276 ph/(m² ns sr) SEE SIMON MACKOVJAK presentation**
 - Remember that situation is different for other positions – background could by higher.
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- [1] Ch. Leinert et al, Astron. Astrophys. Suppl. Ser. 127, 1-99 (1998)
 - [2] D. J. Strickland et al, J. Quant. Spect. Rad. Transfer, 62, 689, 1999.
 - [3] B. Mayer and A. Kylling, Technical note: The libRadtran software package for radiative transfer calculations - description and examples of use, doi:10.5194/acp-5-1855-2005

Sources of background for balloon

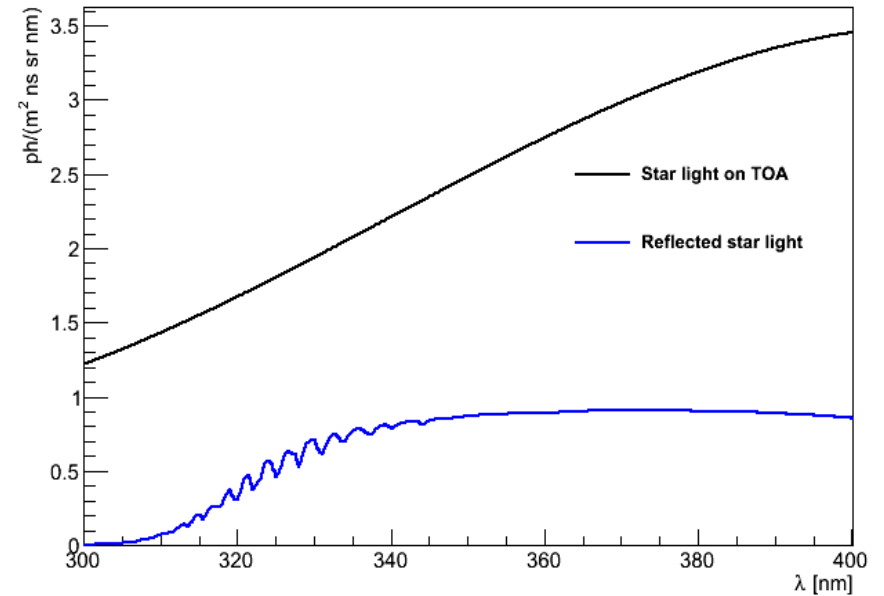
- Integrated star light

- Star light spectrum – fit of spectrum from [1]



[1] Ch. Leinert et al, Astron. Astrophys. Suppl. Ser. 127, 1-99 (1998)

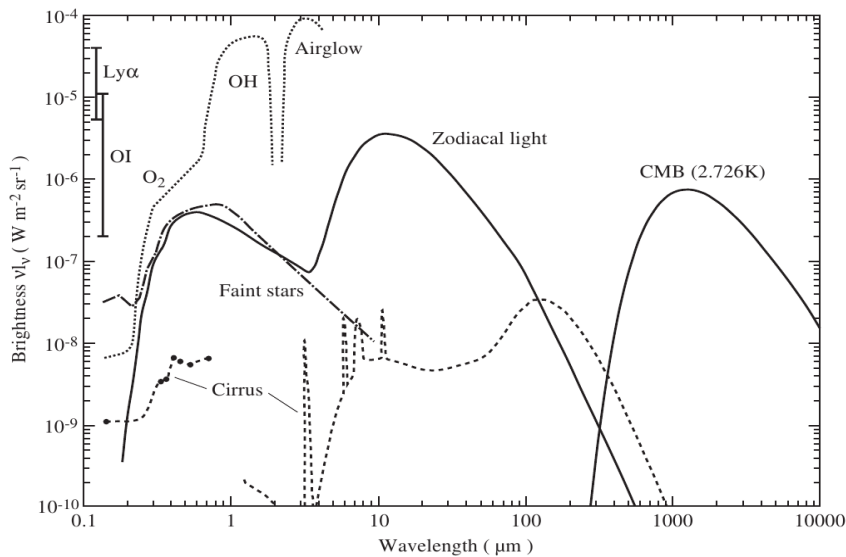
Star light spectrum



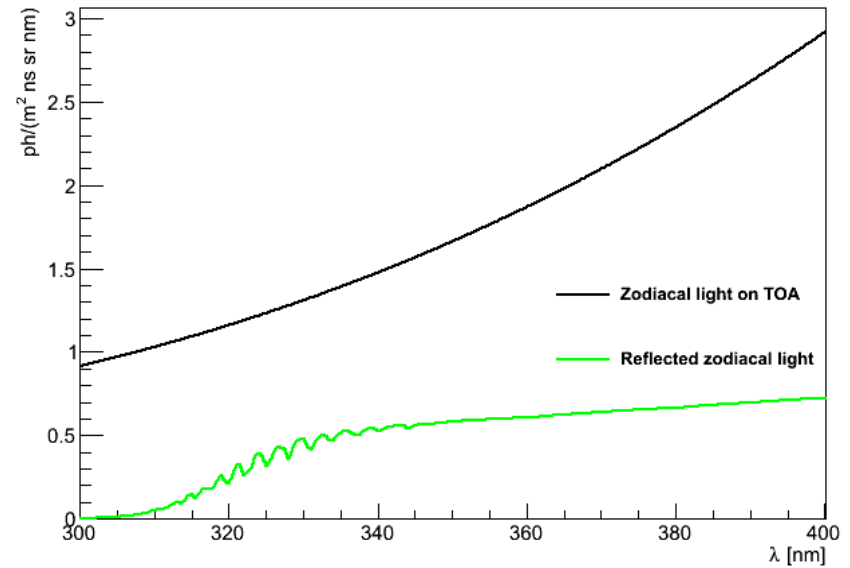
- Integral of TOA spectrum = 243 ph/($\text{m}^2 \text{ns sr}$)
- Integral of reflected spectrum at altitude 40 km = 68 ph/($\text{m}^2 \text{ns sr}$)

Sources of background for balloon - Zodiacal light

- Zodiacal light spectrum – fit of spectrum from [1]



Zodiacal light spectrum



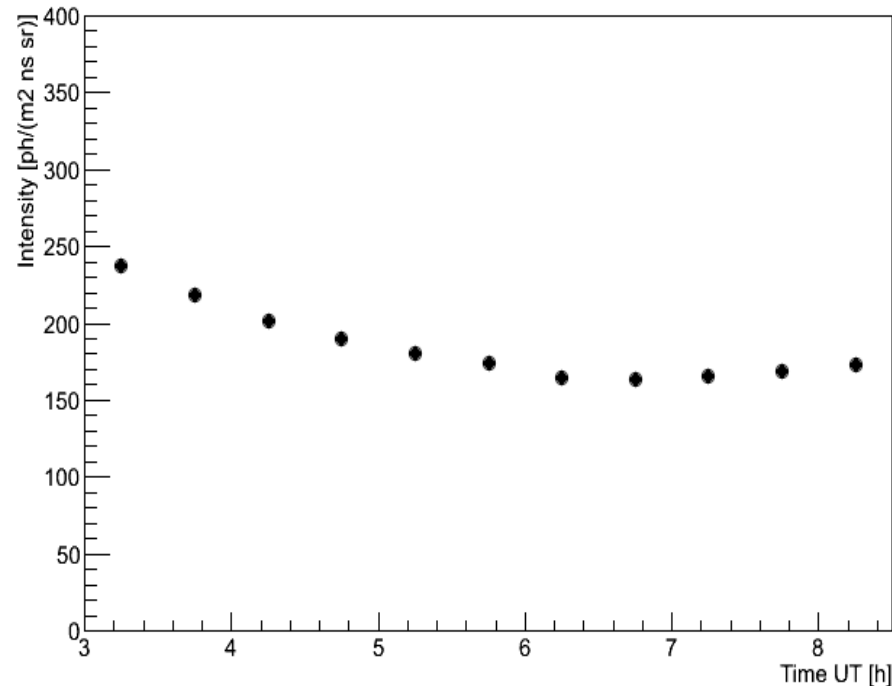
[1] Ch. Leinert et al, Astron. Astrophys. Suppl. Ser. 127, 1-99 (1998)

- Integral of TOA spectrum = 175 ph/($m^2 ns sr$)
- Integral of reflected spectrum at altitude 40 km = 48 ph/($m^2 ns sr$)

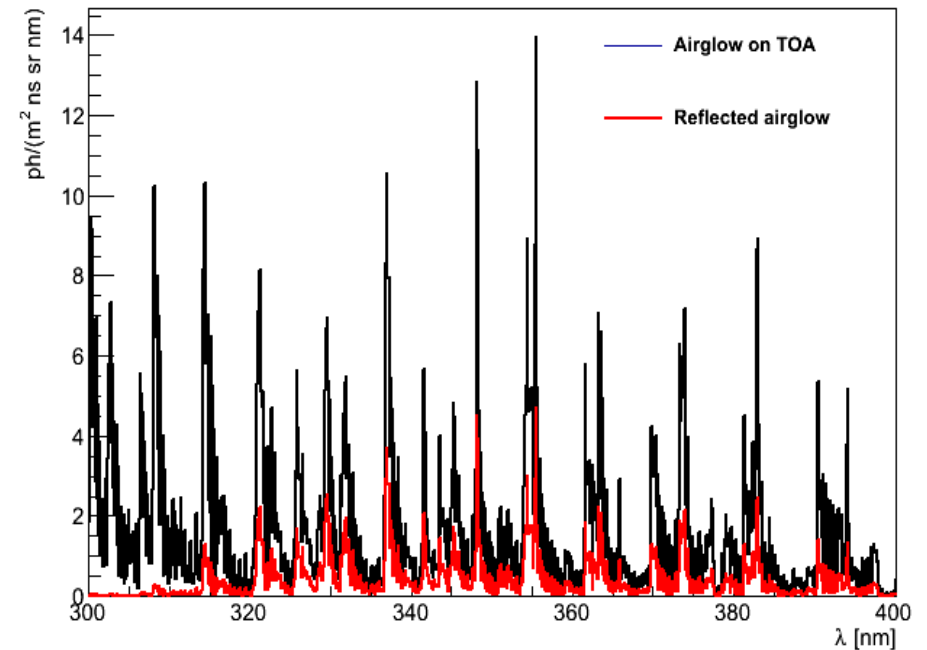
Sources of background for balloon -Airglow

- Provided by AURIC model [1]
- Spectrum at 5:30 UT

Airglow production



Airglow spectrum



- [1] D. J. Strickland et al, J. Quant. Spect. Rad. Transfer, 62, 689, 1999.
- Integral of TOA spectrum = 175 ph/(m² ns sr)
- Integral of reflected spectrum = 43 ph/(m² ns sr)

Overall background

Model (5:30 UT):

- Overall downward intensity:
 - Stars = 243 ph/(m² ns sr)
 - Zodiacal = 175 ph/(m² ns sr)
 - Airglow spectrum = 175 ph/(m² ns sr)
 - Overall = 593 ph/(m² ns sr)
- Overall reflected intensity (40km):
 - Stars = 68 ph/(m² ns sr)
 - Zodiacal = 48 ph/(m² ns sr)
 - Airglow = 43 ph/(m² ns sr)
 - Overall = 159 ph/(m² ns sr)

Data:

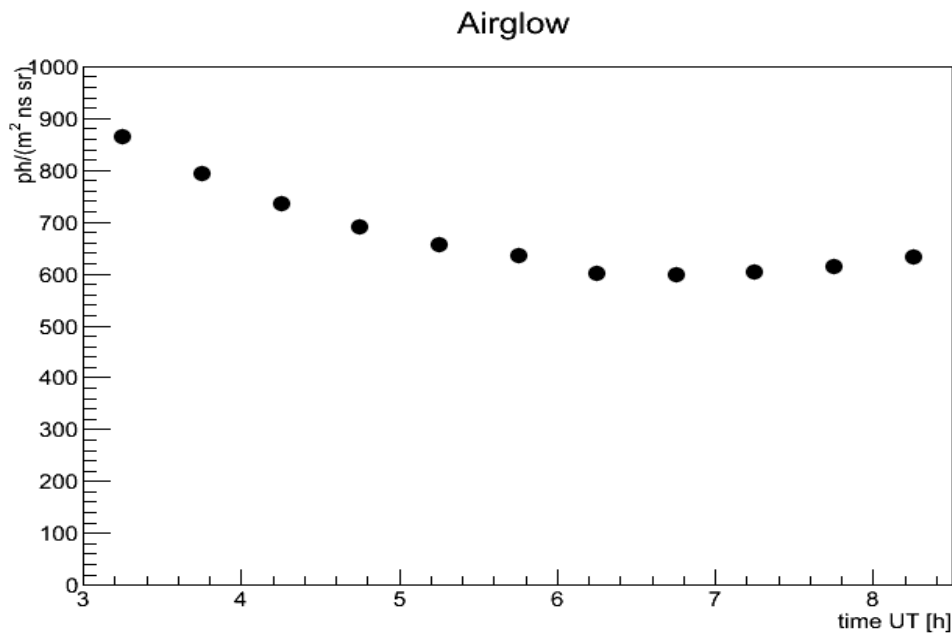
- Preliminary value of background from EUSO-balloon data for no cloud condition at ~ 5:30 UT is ~ **276 ph/(m² ns sr)** SEE SIMON MACKOVJAK presentation
- For conversion of intensity units:

$$1[\text{pe px}^{-1} \text{GTU}^{-1}] * \left(\frac{N_p}{QE OE A t} \right) = 437.64 [\text{ph m}^{-2} \text{sr}^{-1} \text{ns}^{-1}]$$

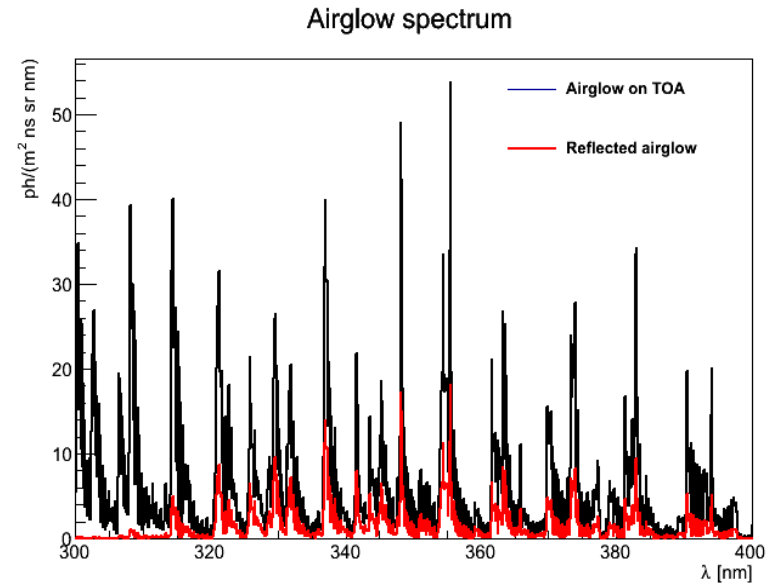
- QE = 0.2
- OE = 0.24
- N_p = 2304
- t = 2500 ns
- A = 144/3282 m² sr

Corrected airglow

- Main parameter of airglow production is atomic oxygen density vertical distribution.
- If we increase it by 100%, we reach agreement with data
- It is too much, something is wrong on model or data side.



- Spectrum at 5:30 UT



- Integral of TOA spectrum = $636 \text{ ph}/(\text{m}^2 \text{ ns sr})$
- Integral of reflected spectrum = $157 \text{ ph}/(\text{m}^2 \text{ ns sr})$
- Overall intensity (airglow + star + zodiacal) = $273 \text{ ph}/(\text{m}^2 \text{ ns sr})$

Possible sources of discrepancy

Model:

- Too many parameters in each model (AURIC, libRadtran)
- Wrong Integrated star and Zodiacal light spectrum
- We need to recheck it all.

Data:

- Are we sure that all factors are included?
- Is structure factor 0.76 included?

$$1[\text{pe px}^{-1}\text{GTU}^{-1}] * \left(\frac{N_p}{QE OE At}\right) = 437.64[\text{ph m}^{-2}\text{sr}^{-1}\text{ns}^{-1}]$$

Part II: Clouds

- Two cloud parameters in libRADtran
 - Effective droplet radius r_e or EDR [μm]
 - Liquid water content w or LWC [g/m^3]

$$w = \int \frac{4}{3} \rho_w \pi r^3 n(r) dr,$$

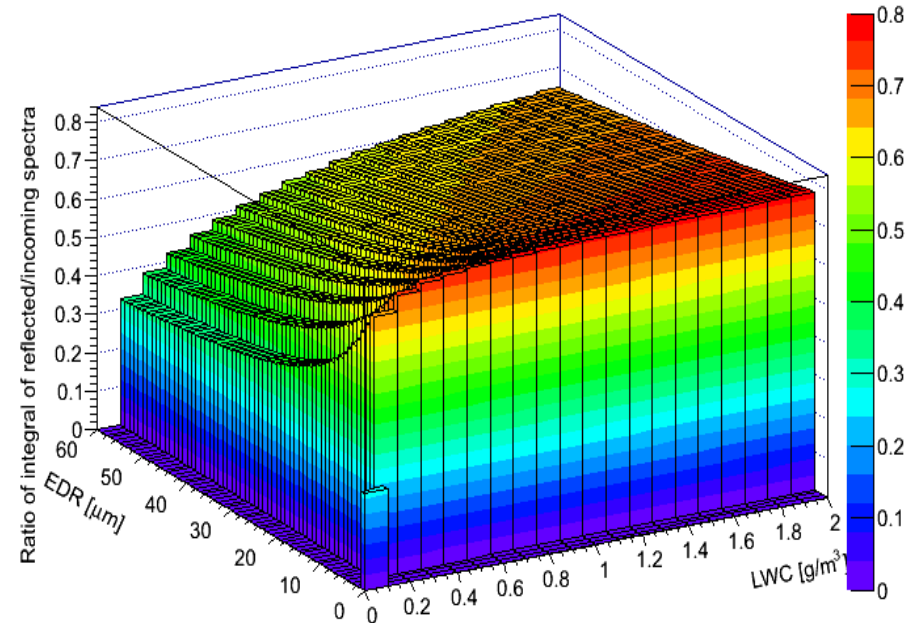
$$r_e = \frac{\int \pi r^3 n(r) dr}{\int \pi r^2 n(r) dr},$$

- r – droplet radius
- ρ_w – water density
- $n(r)$ – droplet size distribution
- We simply try all combinations of EDR (in range 3 – 60 μm) and LWC (0.1 – 2) parameters to obtain first overlook
- Probably not all combinations are physical – we need to find some physical $n(r)$ distributions

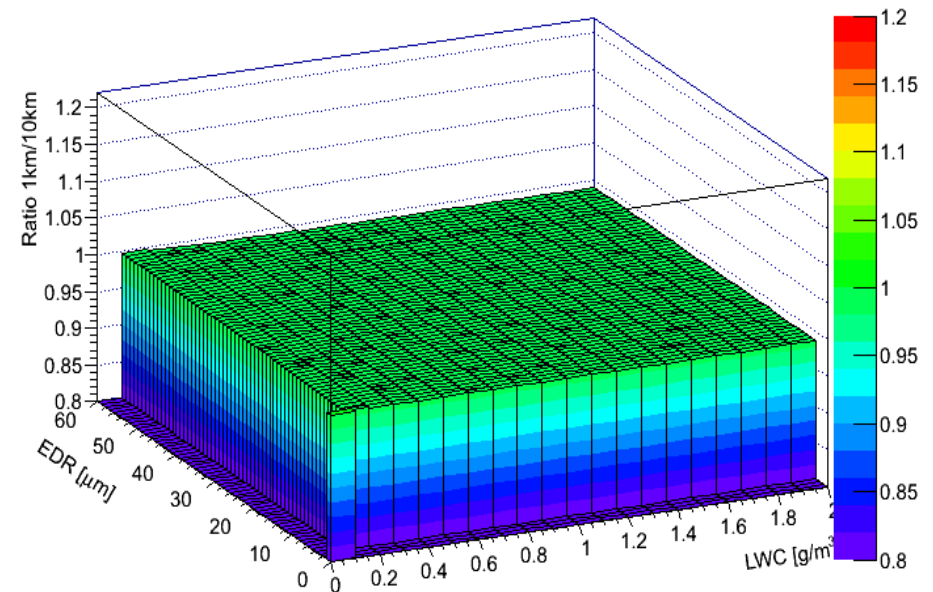
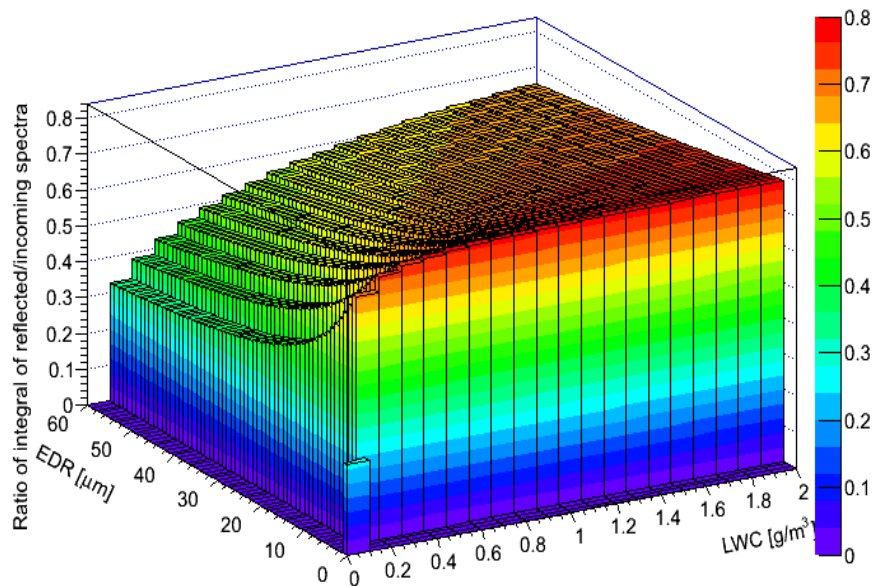
Clouds

- Observer altitude 40 km
- Surface albedo 0
- Cloud altitude 1km and 10km
- Cloud thickens 1km – uniform cloud
- Each bin contain ratio of integral of reflected and incoming spectra.
- Bin 0,0 – value for cloudless situation
- High LWC and small EDR have bigger reflection (~2.6 times) than low LWC and big EDR
- Altitude of cloud change situation in order of few percent

cloud altitude 1km, observer altitude 40 km



cloud altitude 10km, observer altitude 40km



Conclusions

- Model predicted intensity:
 - Standard atomic oxygen density: 159 ph/(m² ns sr)
 - Atomic oxygen density increase 100%: 273 ph/(m² ns sr)
 - It is too much from model point of view, direct airglow radiation increase radically
- Data intensity: 276 ph/(m² ns sr)
- Need some time to understand this

Clouds:

- Altitude of cloud doesn't play significant role (only few percent effect) for reflected intensity at given conditions
- We need to find which combinations of cloud parameters fit real cloud effects
 - In general high LWC and small EDR have bigger (~2.6) reflection than low LWC and big EDR